

Fig. 2: A tephra layer in a half-cut ice core sample. There are a few tephra layers that were not found in the previous 2503 m deep ice core from Dome Fuji.

Properties of the ice core

In situ ECM (Electric Conductivity Measurement) and optical stratigraphic observations will be performed at Dome Fuji. The detailed analysis will be conducted in Japan, and we expect to determine the approximate age of the core by detecting glacial cycles and signals of the Brunhes-Matuyama magnetic reversal at around 780 ka, and by using a glaciological flow-model. It is difficult to say at present but we expect that the core may be 800 ka old, as is the case for the Dome C deep ice core, or even older. This expected age is suggested from the fact that the geothermal heat flux in Dronning Maud Land where Dome Fuji is located is probably smaller than that of the eastern part of East Antarctica where Vostok and Dome C are located. We discovered new

tephra layers (Fig. 3) that were not observed in the first 2503 m deep ice core.

Outlook

One half of the ice core will be kept in a storage trench at Dome Fuji and the other half will be brought back to Japan by our Antarctic research vessel in April. These ice cores will be used for various

analyses at the National Institute of Polar Research, Hokkaido University, and other institutions in Japan. Our main research objectives are

- 1) To reconstruct the global climate and environmental changes that occurred during the past 800 ka to 1 million years.
- 2) To determine the impacts of the geomagnetic field reversal on the global climate and the environment around 780 ka BP.
- 3) To clarify the relationship between the paleo-solar activity and the global climate.
- 4) To clarify the relationship between biological evolution of the ice core microbes and environmental changes.

Furthermore, if we can obtain samples of the bedrock, it may be possible to get information on the formative period of the Antarctic ice sheet in the Tertiary.

REFERENCES

Watanabe, O., Jouzel, J., Johnsen, S., Parrenin, F., Shoji, H., and Yoshida, N., 2003: Homogeneous climate variability across East Antarctica over the past three glacial cycles, *Nature*, **422**: 509-512.

Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N.I., Barnola, J.-M., Basile, I., Bender, M., Chappellaz, J., Davis, M., Delaygue, G., Masson-Delmotte, V., Kotlyakov, V.M., Legrand, M., Lipenkov, V.Y., Lorius, C., Pepin, L., Ritz, C., Saltzman, E., and Stievenard M., 1999: Climate and atmospheric history of the past 420,000 years from the Vostok ice core, Antarctica, *Nature*, **399**: 429-436.

Project facts

Project: Dome Fuji Ice Core
Contact: Yoshiyuki Fujii, fujii@nipr.ac.jp
Participants: Numerous scientists from the National Polar Institute, Tokyo, and several Japanese Universities.
Funding: Japanese National Funding
Where: Eastern Dronning Maud Land, East Antarctica
When: 2004 ongoing
What: Deep ice cores to bedrock; Multi-parameter analysis

Vostok Ice Core project

VLADIMIR LIPENKOV¹ ON BEHALF OF THE VOSTOK PROJECT MEMBERS

¹Arctic and Antarctic Research Institute (AARI), St. Petersburg, Russia; lipenkov@aari.nw.ru

Introduction

Over the last 25 years, the three deep ice cores 3G, 4G and 5G, drilled at the Russian Vostok Station (Fig.1), have provided a wealth of information about past climate and environmental changes. At this site in East Antarctica, the ice thickness is 3750 m and the snow accumulation rate is only 2.1 cm of water equivalent per year. This provides the unique opportunity to

obtain a long climatic record with relatively high time-resolution. In January 1998, the collaborative project between Russia, France and the United States to drill the 5G hole at Vostok yielded the longest ice core ever recovered, reaching a depth of 3623 m. The drilling stopped 130 m above Lake Vostok, a deep subglacial water body that extends over a large area below the ice sheet. The 5G ice core can

be separated into three distinct sections. The upper 3310 m of the core are characterized by an undisturbed sequence of ice layers. The analysis of this section of the core resulted in the first record of Antarctic ice extending through four climate cycles back to 420 kyr BP (Petit et al., 1999). Between 3310 and 3539 m, there are indications of ice-flow anomalies that could have altered the original stratig-





Fig. 1: Drill units at the Russian Vostok Station in East Antarctica.

raphy of glacial ice. Finally, below 3539 m, the core consists of ice re-frozen from lake water. Using an appropriate correction of the ice stratigraphy for flow disturbance in the 3320-3345 m interval, it has recently been possible to extend the Vostok ice record farther back to 440 kyr BP, which implies full

coverage of marine isotope stage (MIS) 11 (Raynaud et al., 2005).

The 440 kyr climatic record from the Vostok ice core

Figure 2 shows a series of selected Vostok records covering the last 440 kyr (adopted from Petit et al., 1999; Delmotte et al., 2004 and Raynaud

et al., 2005). The records indicate that the climate on our planet during this time has always been in a state of change but with atmospheric and climate properties oscillating between stable bounds. Most of the climate variability during glacial-interglacial changes occurs with periodicities corresponding to that of the precession, obliquity and eccentricity of the Earth's orbit, with a larger concentration of variance in the 100-kyr band. The overall amplitude of the glacial-interglacial temperature change is 12°C at the ice-sheet surface, as judged from the Vostok isotope record.

The "sawtooth" pattern of the Vostok isotopic temperature record roughly mimics the sea level (reversed global ice volume) changes deduced from marine sediment studies. The much higher dust concentration during full glacial periods than during interglacials is interpreted as indicating more extensive deserts and continental areas (low sea

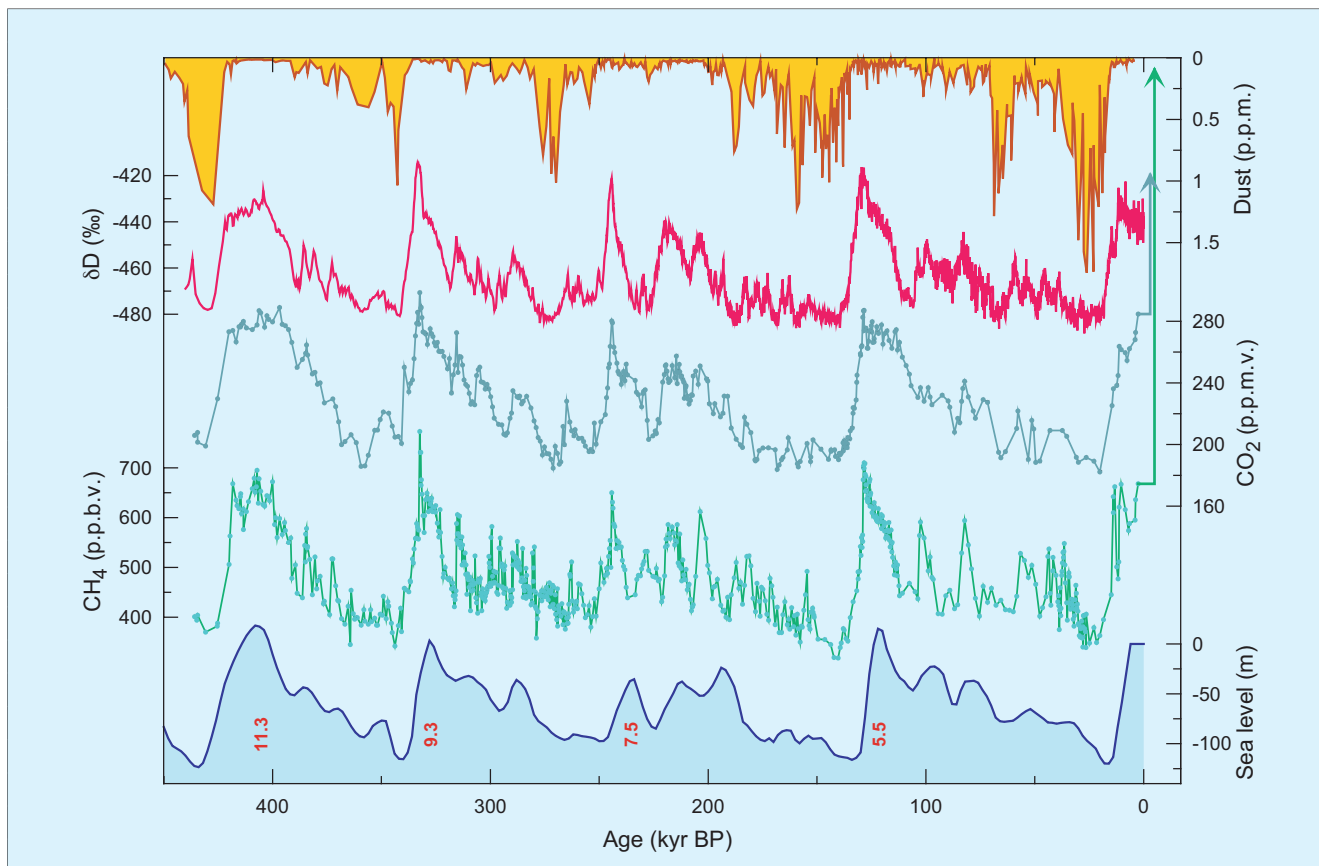


Fig. 2: Vostok time series and sea level derived by calibrating the marine $\delta^{18}\text{O}$ record (Bassinot et al., 1994). The dust record provides information on aerosols of continental origin. The deuterium content of the ice δD is taken as a proxy for Antarctic temperature; the warmest periods of the past correspond to MIS 5.5, 7.5, 9.3 and 11.3. The atmospheric concentrations of the two greenhouse gases (CO_2 and CH_4) are measured on the air enclosed in the ice. The rapid anthropogenic rise in concentrations of CO_2 and CH_4 since the 1850's, up to present-day levels of 360 ppmv and 1750 ppbv, respectively, are indicated by arrows. The Vostok data (Petit et al., 1999; Delmotte et al., 2004; Raynaud et al., 2005) are plotted on the EDC2 timescale after Raynaud et al., 2005.

level), more intense surface winds in the source regions, and/or more efficient meridional transport at the times of glacial maxima (Petit et al., 1999). The close linkage between climate and CO₂ and CH₄ concentrations documented in the Vostok core throughout the four climatic cycles supports the role of the greenhouse gases as amplifiers of initial orbital forcing.

The sequence of events during glacial terminations suggests that increases in Antarctic air temperature and atmospheric greenhouse gas concentrations lead sea level rise (Northern Hemisphere deglaciation) by a few thousands of years (Petit, et al., 1999, Sowers et al., 1991, Shackleton, 2000). It has also been shown that CH₄ (Delmotte et al., 2004) and CO₂ (Caillon et al., 2003; Fischer et al., 1999; Petit, et al., 1999; Pépin, et al., 2001) concentration increases lagged Antarctic warming by several hundreds of years. This natural scenario is, however, different from the present situation, where rise of greenhouse gases, as a result of anthropogenic activities, has been imposed first. The reconstructed record for MIS 11 indicates that concentrations of CO₂ and CH₄ throughout this 30-kyr-long interglacial period, which is considered an "orbital" analog for the Holocene, were close to the pre-industrial levels and lower than during MIS 9.3 (Raynaud et al., 2005). Thus, the present-day increase in concentration of these greenhouse gases in the Earth's atmosphere seems to have been unprecedented during the past 440 kyr.

Project facts

Project: Vostok ice core project

Contact: Vladimir Lipenkov,
lipenkov@aari.nw.ru

Participants: Institutes and laboratories in Russia (AARI of Roshydromet; St. Petersburg Mining Institute; Institute of Microbiology, RAS; Kazan State University; Institute of Geography, RAS), in France (LGGE, CNRS, Université Joseph Fourier, Grenoble; Laboratoire des Sciences du Climat et de l'Environnement, CNRS-CEA, Saclay) and in the United States (laboratories funded by NSF).

Funding: The Vostok venture was made possible by the logistical support of Russian Antarctic Expedition (RAE), Institute Français pour la Recherche et la Technologie Polaires (IFRTP) and the Office of Polar Programs (NSF). The development of drilling technology and the ice-core studies were funded by national contributions from Russia, France and the United States.

Where: East Antarctica (78°28'S, 106°48'E, altitude 3488 m asl, ice thickness 3750 m)

When: 3 main deep holes (3G, 4G, 5G) were drilled from 1980 to 1998. The ice core analyses are in progress.

What: 3 ice cores reaching depths of 2202 m, 2546 m and 3623 m. High-resolution records of atmospheric composition and climate covering the last 440 kyr.

Database: www.ncdc.noaa.gov/paleo/icecore/antarctica/vostok/vostok.html

Ongoing ice core studies

Studies of the Vostok ice are now focused on the deepest section of the core, below 3310 m. The extrapolation of the Vostok timescale to greater depths, made with the aid of an ice-sheet flow model, shows that the age of glacial ice just above its contact with accretion lake ice (at a depth of 3530 m) may reach ~2000 kyr (Salamatin et al., 2004). We hope

that the joint effort of scientists working in the fields of paleoclimate records and physical properties of ice will allow us to decipher information on some of the earlier climatic cycles archived in this section of the Vostok core. Meanwhile, comprehensive analysis of the now available accretion ice from depths below 3539 m is expected to provide a clue to understanding the extraordinary environment of subglacial Lake Vostok.

REFERENCES

- Caillon, N., Severinghaus, J., Jouzel, J., Barnola, J.M., Kang, J. and Lipenkov, V.Ya., 2003: Timing of Atmospheric CO₂ and Antarctic temperature changes across Termination III. *Science*, **299**: 1728-1731.
- Masson-Delmotte, V., Chappellaz, J., Brook, E., Yiou, P., Barnola, J.M., Goujon, C., Raynaud, D. and Lipenkov, V.Ya., 2004: Atmospheric methane during the last four glacial-interglacial cycles: Rapid changes and their link with Antarctic temperature. *J. Geophys. Res.*, **109**: D12104, doi: 10.1029/2003JD004417.
- Fischer, H., Wahlen, M., Smith, J., Mastroianni, D. and Deck, B., 1999: Ice core records of atmospheric CO₂ around the last three glacial terminations. *Science*, **283**: 1712-1714.
- Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N.I., Barnola, J.-M., Basile, I., Bender, M., Chappellaz, J., Davis, M., Delaygue, G., Masson-Delmotte, V., Kotlyakov, V.M., Legrand, M., Lipenkov, V.Y., Lorius, C., Pepin, L., Ritz, C., Saltzman, E., and Stievenard M., 1999: Climate and atmospheric history of the past 420,000 years from the Vostok ice core, Antarctica, *Nature*, **399**: 429-436.
- Raynaud, D., Barnola, J.M., Souchez, R., Lorrain, R., Petit, J.R., Duval, P. and Lipenkov, V.Ya., 2005: The record for marine isotopic stage 11. *Nature*, **436**: 39-40.
- Salamatin, A.N., Tsyganova, E.A., Lipenkov, V.Ya. and Petit, J.R., 2004: Vostok (Antarctica) ice-core time-scale from datings of different origins. *Ann. Glaciol.*, **39**: 283-292.

For full references please consult:
www.pages-igbp.org/products/newsletters/ref2006_1.html



European Project for Ice Coring in Antarctica (EPICA)

ERIC WOLFF¹ ON BEHALF OF THE EPICA COMMUNITY

¹British Antarctic Survey, Cambridge, UK; ewwo@bas.ac.uk

Introduction

The last few hundred thousand years form the context in which we can learn how the Earth System (including its climate) works. Although the geological setting of the Earth was similar to today, very significant changes occurred in the climate, the circulation of the ocean, and the composition of

the atmosphere. By learning about these changes, we can understand the processes that can occur, and that should be represented by models—the same models that will be used to predict future conditions.

Ice cores are particularly powerful because numerous climate responses and forcings (including greenhouse gas concentrations) are

recorded in the same core. In early Antarctic cores covering complete glacial-interglacial cycles (notably Vostok and Dome Fuji), we were able to see the close connection between climate and greenhouse gas concentrations that prevailed through four glacial-interglacial cycles over the last 400 kyr (Petit et al., 1999, Watanabe et al., 2003).